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An assessment of field-portable X-ray fluorescence analysis for nickel and iron in laterite ore (New Caledonia)

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ABSTRACT: The aim of this work was to determine how in-situ portable XRF (pXRF) analysis of ores (in this study, nickel and iron) may be consistent with data obtained from well-prepared samples analysed by conventional XRF methods, and to investigate potential sources of error in field measurements due to different chemical composition, moisture content and particle size. For this study, 201 samples of saprolite and laterite, sieved at 125 μm, were selected for their Ni and Fe contents (from 0.119 to 5.225% Ni and from 4.7 to 37.2% Fe for saprolite samples and from 0.38 to 2.94% Ni and from 21.5 to 55.55% Fe for laterite samples), compressed to make powder pellets and used to calibrate the pXRF. It is shown that for these two elements, pXRF can correlate extremely well (r² > 0.98, relative standard error <7%) with conventional laboratory results when samples are suitably prepared (dried, pulverized and pelletized). Laboratory analysis of a few artificially wetted samples shows that both Ni and Fe estimated concentrations are directly proportional to moisture content. In the field, two experiments were conducted, one on an air-dried drill-core and the second directly on the pit face. The results indicate that because of sample heterogeneity, Ni and Fe contents are better estimated in the laterite layer than in the saprolite layer. Despite the many challenges associated with field conditions, pXRF can provide useful benefits for a number of applications in the mining sector.

KEYWORDS: portable XRF, nickel ore, iron, in-situ analysis, mining, moisture, granulometry

During all phases of the mining process, from exploratory screening to the quality control of final products, the mining industry has a wide array of testing and analysis needs. Elemental and mineral analyses are usually performed in the laboratory. The actual workload, which involves several steps such as sampling, drying, grinding and analysis, takes time (from several hours to days). Having access to metal content and concentration data in real time, directly in the field, would help geologists in making decisions and save time, effort and also money. In recent years, portable analytical tools have emerged. They can be classified into two categories depending on the analytical information provided, molecular or atomic. The first, based on infrared or Raman technologies, is dedicated to mineral identification (e.g. the Agilent FlexScan, Thermo Scientific TruDefender, Intevac RockHound), whilst the second, based on XRF (e.g. Thermo Scientific XL3t, Olympus Innov-X) or LIBS (Ivea MobiLIBS) technologies (Fortes & Laserna 2010), allows multi-elemental chemical analysis to be performed.

Field-Portable X-ray fluorescence (pXRF) is currently used in many different disciplines, for example: environmental remediation (Kalnicky & Singhvi 2001; Laperche et al. 2008; Higueras et al. 2012), recycling, archaeological applications (Cesareo et al. 2009), health and geology/mining (Houlahan et al. 2003). Thanks to the commercialisation of miniature X-ray tubes and the development of cost-effective Silicon Drift detector (SDD) technology, pXRF technology is becoming more affordable. Because the technique is non-destructive, multi-elemental, fast and cost-effective and can be directly applied to certain types of samples without any preparation, it is ideally suited for a number of applications. A good example is archaeology, where pXRF allows in-situ analysis of historical objects that cannot be easily moved to the laboratory. In the mining context, one would expect a similar revolution but, at least in New Caledonia, miners are cautious because analysis is often inaccurate and imprecise on natural samples when no preparation is envisaged or possible in the field. Whatever the XRF device used, the sample (standards as well as unknowns) must always be prepared with strictly the same protocol. Chemical and physical matrix effects (Markowicz 2008) are sources of error unless methods to correct for them are put in place.

Although processors and embedded signal processing devices are becoming more sophisticated and efficient, no instrument can fully compensate for all chemical and physical effects. All operators should keep this limitation in mind and analyse results with a critical eye (e.g. measurement uncertainty), especially when analysing raw samples in the field. Therefore, knowledge of capabilities and limitations of the instrument and sample features is critical. Portable XRF
instruments are more than a simple point-and-shoot technology and this knowledge can only be obtained if time has been devoted to methodology and data quality objectives.

The aim of this work is to study potential error sources in field measurements due to different chemical composition, moisture content and particle size. Because measurements by pXRF are influenced by these factors, it is important to investigate their impact and determine the best adjustment factors in order to achieve results comparable to laboratory analyses (e.g. ICP-MS, ICP-AES).

STUDY SITE

New Caledonia, lying just north of the Tropic of Capricorn, is an island group located in the SW Pacific, 2000 km east of Australia. It comprises one main mountainous island, known as the ‘Grande Terre’, surrounded by a barrier reef, and several smaller islands. New Caledonia is the world’s third largest producer of nickel ore, after Russia and Canada, and yields c. 25% of the world’s lateritic nickel reserves. With increasing nickel mining by both private and public companies, the nickel industry is becoming even more important to New Caledonia’s economy than in the past.

New Caledonia’s nickel resources come from the tropical weathering of ultramafic rocks (peridotites) that cover about one third of the Grande Terre. These rocks were tectonically emplaced (obducted) over New Caledonia during the Late Eocene (c. 34 Ma; Cluzel et al. 2001) and are referred to as ‘Peridotite Nappe’. Peridotites mainly outcrop in the south-eastern part of the island (Massif du Sud) and in smaller massifs distributed along the west coast (Fig. 1).

Four main horizons are recognized in weathering profiles developed upon peridotites (Fig. 2), with, from the base to the top: (1) coarse saprolite (S1) developed over the bed-rock (P) where Ni is concentrated in residual and neoformed hydrous silicates (nickeliferous serpentine, talc and smectite, referred to as garnierite); (2) fine-grained yellow laterite (L1) mainly composed of Ni-bearing Fe oxy-hydroxides; (3) red laterite (L2), a granular material mainly composed of Ni-poor Fe oxy-hydroxides; and (4) ferricrete on top (F1 and F2). On some profiles, a thin discontinuous layer of granular cobalt (Co) ore (asbolane; S2) often enriched in chromium (Cr) and manganese (Mn) oxides, appears on top of the coarse saprolite horizon. This sub-layer is named fine saprolite.

The typical composition of a theoretical profile is given in Table 1 (from Troly et al. 1979). It shows that Fe content increases along the weathering profile from base to top while silica and magnesia decrease. Silica and Fe are major chemical elements in saprolites and laterites, respectively.

Nickel ore in New Caledonia is worked in open pit mines and c. 9 Mt of wet ore was produced in 2010, which consists of 70% saprolitic and 30% lateritic ore. Two different processes have been developed for extracting Ni from saprolite (silicate ore), or laterite (oxidised ore). Low-Fe saprolite ore is processed by pyrometallurgy to be transformed into ferro-nickel and matte in a smelter plant, Société Le Nickel (SLN) and Koniambo Nickel SAS (KNS). Laterite processing is based on hydrometallurgy (Vale NC), which allows Ni to be extracted by acid leaching (sulphuric acid).

MATERIALS AND METHODS

Different types of samples were used for this study, such as well-prepared samples for the calibration of the pXRF in the laboratory, core samples to test the calibration on-site and fresh samples analysed directly on the working face. The chemical analyses of all the samples, used in this study to be compared to our pXRF data, were performed by XRF spectrometry (PW 1606, Panalytical) on glass beads by borate fusion (1 g of sample for 9 g of Li₂B₄O₇).

pXRF

All pXRF analyses were performed using the Niton XL3t-900 GOLDD model. This energy dispersive pXRF is fitted with a 50 kV X-ray tube (max. 50 kV, 100 μA, 2 W) with an Ag anode target excitation source and a Large Drift Detector (LDD). The analysed spot had an average diameter of 8 mm. As part of
FPXRF Analyzer for Ni and Fe content of Laterite Soils

The samples (c. 3 g) were then placed in a press to produce pellets. In order to evaluate the minimum thickness of the pellets, the penetration depth of X-rays has been calculated (Potts et al. 1997) for different rocks. As can be seen from Table 2, the critical penetration depth is limited to a few hundred µm. In this study, all samples were prepared so that their thickness exceeds 2 mm (average 3 mm).

pXRF calibration. All the pellets were analysed with the pXRF and the results used to determine two sets of XRF calibration factors for Ni and Fe elements, one for the saprolites and the other for the laterites. The calibration was performed on pulverized and pelletized samples. Since the number of samples is large, each sample was measured once with 30 s duration. In the particular case of laterites, two additional analyses were performed by placing a Kapton or a PVC foil between the instrument head and the sample. In both cases, the foils had a thickness of 110 µm.

Temporal and spatial standard deviation. In order to evaluate the temporal and spatial standard deviation of raw concentrations, four pelletized samples were selected. Temporal dispersion is a measure of instrument precision while spatial dispersion is a measure of both instrument precision and sample heterogeneity. For each sample, a total of 18 pXRF analyses were carried out on 6 different zones (p=6) on both sides of the pellet. Each zone was then measured 3 times (q=3), leaving the head instrument at exactly the same place. $C_{ji}$ is the $j$th measure of the raw concentration in the $i$th position. In this way, both temporal and spatial standard deviation can be calculated.

For each position $p_i$, the local average concentration $m_i$ ($i=1…p$) can be calculated from the $q$ measurements:

$$m_i = \frac{1}{q} \sum_{j=1}^{q} C_{ji}$$

The local sample standard deviation $s_i$ is given by:

$$s_i = \sqrt{\frac{1}{q-1} \sum_{j=1}^{q} (C_{ji} - m_i)^2}$$

and the global average concentration of the sample by:

$$m = \frac{1}{p} \sum_{i=1}^{p} m_i$$

We define the temporal standard deviation $s_t$ as the average of local standard deviations over the $p$ zones and the spatial standard deviation $s_m$ as the standard deviation of local averages. Therefore, the temporal standard deviation is given by:

$$s_t = \frac{1}{p} \sum_{i=1}^{p} s_i$$

and the spatial standard deviation by:

$$s_m = \sqrt{\frac{1}{p-1} \sum_{i=1}^{p} (m_i - m)^2}$$

Laboratory analysis

For this study, 98 samples of saprolite and 103 samples of laterite, sieved at 125 µm, were retrieved from the sample archives of mining companies. In addition, borate fusion XRF reference values were also delivered for all samples. The choice of these samples was based on Ni and Fe contents from 0.119 to 5.225 % Ni and from 4.7 to 37.2 % Fe for saprolite samples and from 0.38 to 2.94 % Ni and from 21.5 to 55.55 % Fe for laterite samples.

Fig. 2. The four horizons of the peridotite weathering profile.

P. Peridotite, $S_1$: Coarse Saprolite, $S_2$: Fine Saprolite, $L_1$: Yellow Laterite, $L_2$: Red Laterite, $F_1$: Granular Iron and $F_2$: Iron Duricrust (Picture from Vale NC).
Grain size analysis. Attempting to measure chemical composition of rocks directly poses the problem of the best sampling, because particle size distribution and vein occurrence can drastically influence results. In order to evaluate the effects of particle size distribution, three samples (one coarse saprolite and two laterites) were sieved in a sieve shaker. Five dry grain fractions (< 80 µm, 80 µm–400 µm, 400 µm–800 µm, 800 µm–1.6 mm and 1.6 mm–4 mm) were isolated, weighed and analysed through a plastic bag by pXRF. Since heterogeneity increases with grain size, five measurements were taken and averaged for the finest fractions (< 400 µm), while 10 measurements were taken for the coarser fractions.

Influence of water content. In order to test the influence of water content on the analysis results, four more samples were collected on working faces: red laterite, yellow laterite, fine saprolite and coarse saprolite. The fresh samples were ground with a mortar and pestle and passed through a 2-mm sieve. This fraction was used instead of the < 125 µm fraction because it was not possible to produce wet pellets with very fine particles mixed with a high water content. Furthermore, this fraction is closer to the field conditions than the < 125 µm fraction. Before producing the pellets, c. 2 g of sample were mixed with different amounts of water. Above a certain quantity of water, the sample/water mixture becomes too liquid to form a pellet. The maximum water content varies, depending on the nature of the sample. In the case of laterites, the water content was up to 35%, whereas for fine and coarse saprolites, it was up to 44 and 66%, respectively. For each sample, 3 measurements were taken on different areas of the pellet. The moisture $H_w$ is expressed as a percentage of the sample’s wet weight:

$$H_w = \frac{m_w - m_d}{m_w} \times 100$$

where $m_w$ and $m_d$ are the masses of the wet and dry sample, respectively.

In-situ analysis

Drill-core. The 42-m long drill-core was stored in 1-m length labelled sample boxes. All the cores were first logged by a geologist and properties used in the core description recorded by intervals. An interval represents a change in the weathering profile but its maximum length was limited to 1 m. On average, interval length was 0.7 m (from 0.2 to 1 m). The core drilling was split longitudinally and half of the split core was sent to the laboratory to be analysed by the lithium metaborate fusion method (‘ME-ICP93’) and ICP-AES spectrometry.

At the time when the measurements were taken on site, the core was air-dried. The pXRF analyses were performed in-situ, by placing the analyser probe directly on the planar surface of the remaining half-core. Two experiments were undertaken, one week apart. During the first experiment, a single measurement of 15-s duration was made in the middle of each interval. During the second experiment, five analyses were performed using a regular distance between the measurements along the interval. The single measurement and the average of the five measurements were then compared to the analyses performed by fusion ICP-AES on the other half-core.

Pit fan. The pXRF analyses were performed in-situ, by placing the analyser probe directly on the surface on 57 fresh samples. Twenty-seven samples were analysed directly on the working faces (3 to 5 measurements were taken, depending on the heterogeneity of the sample) and 30 samples, about one cubic decimetre, were collected from the working face (10 measurements were taken on the different sides of each sample). The 57 samples were identified as 24 red laterites, 16 yellow laterites and 17 fine saprolites. The pit face did not yet reach the coarse saprolite level. After pXRF analysis, the samples were split into two parts, one for XRF analysis and the other for the water content determination in an oven at 105 °C for 24 h.

RESULTS AND DISCUSSION

pXRF analyses in the laboratory

pXRF calibration. Quantitative analysis is based on the ‘Fundamental Parameter Method’ which provides multi-site capabilities. This method uses X-ray theory to mathematically predetermine inter-element coefficients and thus corrects for all matrix effects that affect the intensity emitted by a given element. However, uncertainties in the calculation of coefficients may lead to systematic biases. To obtain reliable results, adjustments based on reference samples, which are representative of the materials under study, may still be needed. Systematic biases can be overcome through a calibration process that consists in finding, for a given element, the relationship between

### Table 1. Typical chemical composition (in %), relative density (g/cm³) and water content (in %) of weathered peridotite in New Caledonia (from Troly et al. 1979). The numbers in the first row refer to the peridotite weathering profile (Fig. 2).

<table>
<thead>
<tr>
<th>Horizon</th>
<th>NiO</th>
<th>Fe₂O₃</th>
<th>SiO₂</th>
<th>MgO</th>
<th>$p_d$</th>
<th>Moisture (%)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Iron Duricrust (F₃)</td>
<td>0.3</td>
<td>75</td>
<td>1</td>
<td>0.1</td>
<td>3</td>
<td>5</td>
</tr>
<tr>
<td>Iron (F₄)</td>
<td>0.9</td>
<td>72</td>
<td>1</td>
<td>0.6</td>
<td>2.2</td>
<td>20</td>
</tr>
<tr>
<td>Red Laterite (L₁)</td>
<td>1.4</td>
<td>70</td>
<td>30</td>
<td>16</td>
<td>1.9</td>
<td>40–50</td>
</tr>
<tr>
<td>Yellow Laterite (L₂)</td>
<td>2.3</td>
<td>32</td>
<td>3</td>
<td>16</td>
<td>1.7</td>
<td>35</td>
</tr>
<tr>
<td>Earthy Saprolite (S₁)</td>
<td>2.3</td>
<td>32</td>
<td>35–39</td>
<td>21–28</td>
<td>1.8–2.4</td>
<td>22–30</td>
</tr>
<tr>
<td>Massive Saprolite (S₂)</td>
<td>2.3</td>
<td>8</td>
<td>44</td>
<td>45</td>
<td>2.6–3.2</td>
<td>5</td>
</tr>
<tr>
<td>Peridotite (Harzburgite) (P)</td>
<td>0.3</td>
<td>8</td>
<td>44</td>
<td>45</td>
<td>2.6–3.2</td>
<td>5</td>
</tr>
</tbody>
</table>

### Table 2. Critical penetration depths $d$ (µm) of different rock types (oxide concentrations, in %, from Pelletier 2003).

<table>
<thead>
<tr>
<th>Rock type</th>
<th>NiO</th>
<th>FeO</th>
<th>CoO</th>
<th>SiO₂</th>
<th>MgO</th>
<th>H₂O</th>
<th>$d_{H}$ (µm)</th>
<th>$d_{C}$ (µm)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Olivine</td>
<td>0.4</td>
<td>8</td>
<td>0.02</td>
<td>41</td>
<td>50</td>
<td>0</td>
<td>190.2</td>
<td>182.5</td>
</tr>
<tr>
<td>Orthopyroxene</td>
<td>0.08</td>
<td>5</td>
<td>0.02</td>
<td>60</td>
<td>35</td>
<td>0</td>
<td>210.5</td>
<td>176.1</td>
</tr>
<tr>
<td>Serpentinite</td>
<td>0.4</td>
<td>7</td>
<td>0</td>
<td>40</td>
<td>39</td>
<td>13</td>
<td>216.3</td>
<td>207.3</td>
</tr>
<tr>
<td>Harzburgite</td>
<td>0.3</td>
<td>7</td>
<td>0</td>
<td>41</td>
<td>42</td>
<td>7</td>
<td>210.6</td>
<td>199.2</td>
</tr>
</tbody>
</table>
pXRF data (raw data \(C_r\)) and XRF, ICP-AES or AAS laboratory data (reference data \(C\)). A linear regression analysis can be performed for this purpose. Linear regression produces the slope \(\alpha\) and the y-intercept \(\beta\) of the regression line. A slope of 1.0 and a y-intercept of 0.0 indicate that pXRF is accurate. If not, the values of \(\alpha\) and \(\beta\) can be entered directly into the analyser. The corrected concentration \(\hat{C}\) is then defined by:

\[
\hat{C} = \alpha \cdot C + \beta
\]

In addition to calibration factors, linear regression analysis produces statistical parameters that can be used to evaluate the level of data quality. One of them is the coefficient of determination, also called r-squared (\(r^2\)), which is the square of correlation of the coefficient \(r\). The coefficient of determination is a measure of the scatter about the regression line and is a measure of the strength of the linear association. It ranges in value from 0.0 (no linear association) to 1.0 (perfect linear association). The standard deviation \(s\), which is a measure of random error of y-values, and Percent Relative Standard Error (RSE\(_n\)) were also calculated. RSE\(_n\) is defined by:

\[
RSE_n = 100 \times \sqrt{\frac{\sum \left(\frac{C - \hat{C}}{C_i}\right)^2}{n - 1}}
\]

where \(n\) is the number of samples, \(C_i\) and \(\hat{C}_i\) are the reference and corrected values of the \(i^{th}\) sample.

Figures 3 and 4 show scatter plots obtained for Ni and Fe, respectively for the saprolite samples \((n = 98)\). The regression line is above the reference line \((y=x)\) for Ni but is under the reference line for Fe. This means that pXRF underestimates Ni but overestimates Fe, probably because of secondary fluorescence (Ni K\(\alpha\) is absorbed by Fe). This secondary fluorescence can also explain why calibration factors are different between the saprolite and laterite layers. For both Ni and Fe, the coefficient of determination is high (> 0.99).

In order to protect the pXRF instrument from water and fine particles, some pXRF users cover the front of the analyser with a Kapton\textsuperscript{®} sticker foil. Insertion of this foil between the sample surface and instrument head compromises light element measurement (Mg, Si) but also attenuates primary X-rays and fluorescence X-rays of heavier elements (Fe, Ni). During the calibration process of laterites, both situations (i.e. with and without Kapton\textsuperscript{®} foil) were evaluated. Figures 5 and 6 show the scatter plots obtained for Ni and Fe, respectively for the laterite samples \((n = 103)\). Insertion of Kapton\textsuperscript{®} foil causes an increase in the slope \(\alpha\) for both elements and so an increase in the attenuation. In all cases, the coefficient of determination is high (> 0.98).

Table 3 gives an overview of the results. For Ni, the standard deviation \(s\) is higher in the saprolite layer than in the laterite layer. It is lower for Fe. This result is consistent with an increase in average concentration \(C_{average}\). The RSE\(_n\) parameter is less than a few percent, except for Ni in the saprolite layer where it reaches 20.77%. However, this value decreases to 7.09% when smaller concentration values (C<0.3) are excluded from the calculation. This problem is not encountered in the laterite layer, since the lower Ni concentration of the samples base is 0.38%.
In conclusion, at least for Ni and Fe, pXRF can correlate extremely well with conventional laboratory measurements when samples are suitably prepared (dried, pulverized and, if appropriate, pelletized). Although Kapton® foil attenuates the X-rays, the correlation with reference data is still high and the standard deviation is of the same order of magnitude for Ni and Fe. Consequently, it is possible to use Kapton® foil to protect the instrument head if necessary. Unbiased results can be obtained by calibration.

**Temporal and spatial standard deviation.** The bar charts shown in Figures 7 (Ni) and 8 (Fe) display the overall average raw concentrations for the four samples. Error bars on the left and on the right represent standard deviations, $s_n$ and $s_m$ respectively. The temporal dispersion is low compared to the spatial dispersion. This means that even on well-prepared samples, heterogeneity is one of the principal factors that affect data quality.

**Particle size chemistry.** The aim was to check how concentrations change from one fraction to another and to determine, if possible, the overall concentration from individual concentrations. The individual concentration is calculated by weighting the measured concentration by its mass fraction. The number of samples is, of course, too small for a trend to emerge, but the idea was to evaluate the possibility of measuring only the finest fraction and using it to estimate the concentration averaged over all grain sizes.

Figures 9 and 10 show, for one of the studied laterites, Ni and Fe raw average concentrations for the different fractions, one standard deviation ($\pm s$) and the minimum and maximum measured values. As expected, data dispersion increases with grain size. For this sample, the distribution of Ni is almost constant but Fe is mostly in the finest fraction.

Table 4 gives the relative mass of the sample retained on each sieve. The finest fraction is by far the most important since it represents more than 90% of the total.

**Fig. 5. Nickel calibration curves – laterite (□ Without film, ○ Kapton®).**

**Fig. 6. Iron calibration curves – laterite (□ Without film, ○ Kapton®).**

| Table 3. Average concentration ($C_a$ in %), slope ($\beta$), y-intercept ($\alpha$) of the regression line, coefficient of determination ($r^2$), standard deviation ($s$) and relative standard error (RSE) of the saprolite and laterite samples |
|---|---|---|---|---|---|---|
| Analyte | $C_a$ | $\alpha$ | $\beta$ | $r^2$ | $s$ | RSE% |
| Saprolite | Ni | 2.3 | 1.1126 | 0.072 | 0.9918 | 0.13 | 20.77(7.09) |
| | Fe | 18.59 | 0.9225 | -0.4084 | 0.9969 | 0.47 | 2.96 |
| Laterite | Ni | 1.27 | 1.2594 | 0.0062 | 0.9947 | 0.04 | 3.68 |
| | Ni (Kapton®) | 1.6694 | 0.0628 | 0.9887 | 0.06 | 5.47 |
| | Ni (PVC) | 1.6083 | 0.0625 | 0.9881 | 0.05 | 4.7 |
| | Fe | 47.6 | 0.9970 | -3.3676 | 0.9870 | 0.87 | 1.84 |
| | Fe (Kapton®) | 1.2979 | -5.2773 | 0.9899 | 0.77 | 1.79 |
| | Fe (PVC) | 1.2956 | -6.8843 | 0.9954 | 0.45 | 1.09 |
calculated. The reference values are 1.36% for Ni and 61.16% for Fe, whilst the calculated values are 1.49% for Ni and 66.35% for Fe. For this sample, 85% and 97% of total Ni and Fe, respectively, are carried by the finest fraction. One solution to deal with the heterogeneity of laterite samples in the field would be to measure only the finest fraction and to leave out zones comprising coarser grains. The non-measured fraction would then have to be taken into account by a statistical compensation. Measurement of a large number of samples would be necessary to make a statistical inference, but since laterite nickel ores contain significant amounts of very fine material (between 60 and 80%), this process could lead to acceptable results.

Corrections for water content. Moisture and grain size are the most problematic parameters to deal with in the mining context. In order to evaluate how moisture affects concentration measurements, four samples were selected, one saprolite and three different laterites. During the experiment, each sample was mixed with an increasing quantity of water, weighed and measured by pXRF. If $H\%$ is defined as the moisture content of the sample, the concentration loss $CL$ of a given analysis, expressed in percent, is given by:

$$CL(\%) = \frac{C_0 - C_{H\%}}{C_0} \times 100$$

where $C_0$ and $C_{H\%}$ are the raw concentrations measured on the dried and wet sample, respectively. The relationship between concentration loss and moisture is shown in Figures 11 (Ni) and 12 (Fe) for the four samples. In both cases, the points are almost aligned on a regression line and the regression analysis gives a high coefficient of variation ($r^2 > 0.94$). It is possible to write:

![Fig. 7. Temporal (left error bar) and spatial (right error bar) standard deviation of Ni concentration.](image1)

![Fig. 8. Temporal (left error bar) and spatial (right error bar) standard deviation of Fe concentration.](image2)

![Fig. 9. Distribution of Ni concentration vs particle size (Boxplot values: minimum, -s, average (C), +s, maximum).](image3)
Because \( \lambda \) is close to 1.0 for Ni while it is close to 0.8 for Fe, Ni is proportionally more attenuated than Fe. Combining the two equations above, the estimated concentration \( \hat{C}_0 \) can then be calculated from:

\[
\hat{C}_0 = C_{w\%} \frac{100}{100 - (\lambda \times H\%)}
\]

Figures 13 and 14 show the concentrations estimated from the above equation for Ni and Fe, respectively. The horizontal lines correspond to the target values \( C_0 \). To make the figures more legible, an offset of +0.5 was applied on coarse saprolite (Fig. 13) and +5 and -5 on fine saprolite and yellow laterite, respectively (Fig. 14). Except for the fine saprolite, which is underestimated whatever the moisture value, the concentrations are well corrected. Since fine saprolite is an intermediate stage between coarse saprolite and yellow laterite, its behaviour may be explained by its chemical and physical nature (heterogeneity of saprolites, porosity and wetting potential of laterites).
second possibility is that the reference value $C_0 (2.50)$ is slightly overestimated (average of estimated concentrations: 2.25).

**In-situ pXRF analyses**

**Drill-cores.** Since the pXRF was only calibrated from XRF reference values, fusion ICP-AES calibration factors were determined directly from the core. Because of matrix and particle size effects, two calibration sets were determined. The first set covers the upper part, named (H1) in Figures 15 and 16, from iron duricrust (F) to yellow laterite (L1) while the second set covers the lower part (H2), from fine saprolite (S2) to the bedrock (P). The calibration factors, with the y-intercept set to zero in all cases, are 1.23 (H1) and 1.15 (H2) for Ni and 0.94 (H1) and 0.85 (H2) for Fe.

The results of this test are graphically represented in Figures 15 (Ni) and 16 (Fe). Except for two points, it can be seen that a good agreement exists between corrected pXRF and fusion ICP-AES reference data in the upper part H1. In the H2 section, data are more dispersed since the particle size is increasing and the material becoming more heterogeneous.

Percent relative standard errors, RSE%, are presented in Table 5. In the calculation, the two odd points have been removed. In the upper part (H1) of the core, the average of five measurements on a section of about one metre provides correct results. In the second part (H2), except for indurated rock...
transitions noted (P) in Figure 15, RSE% is twice for Ni and six times more for Fe compared to the upper part (H₁). In order to obtain similar results, one solution would consist of increasing the number of shots.

Pit face. Fifty-seven fresh samples, from red laterites to fine grained saprolites, were measured directly in the field without preparation. Samples were collected and sent to laboratories in order to obtain both reference concentrations and moisture contents. Nickel concentrations range from 0.78 to 2.55% and Fe concentrations from 26.8 to 56.6%. Moisture ranges from 30 to 50% depending upon the sampling location. As a general rule, moisture is higher in red laterites than in fine-grained saprolites. Concentrations were then corrected as described previously. Figure 17 shows the raw concentrations and moisture-corrected concentrations versus Ni reference data. Because moisture attenuates X-rays, the slope of the regression line is high (1.96) before correction. The coefficient of determination \( r^2 \) is lower (0.77) than on well-prepared dry samples because of particle size chemistry, surface effects, sampling and differences in moisture content from sample to sample. Indeed, the slope of the regression line shows average effects (average moisture) over the sample set while dispersion shows individual effects (moisture of a particular sample). After correction for moisture, \( \lambda=1 \), the slope decreases to 0.94, showing that average moisture is fairly well compensated. The value of \( r^2 \) (0.83) is only very slightly greater than the previous value.

Figure 18 shows the scatter plots of raw and corrected \((\lambda=0.8)\) concentrations for Fe. The slope decreases from 1.29 to 0.96, showing here also that moisture is on average fairly well compensated. The value of \( r^2 \) increases from 0.34 to 0.72, so the gain is more important for Fe than for Ni. There is no obvious reason for this, although thorough analysis of physical phenomena including dilution and matrix effects may help this result to be better understood.

CONCLUSIONS

Portable XRF methodology is considered to be a viable, cost- and time-effective approach for on-site analysis of environmental samples. In the mining context, different situations and objectives may be encountered. If used as a bench-top laboratory unit, pXRF can provide accurate and precise elemental analysis providing that the instrument is properly calibrated for the matrices and elements of interest and that samples are well-prepared (finely ground and dried). This technology is especially suitable for use by small companies because it can be used for laboratory and on-site analysis as well. In the field, however, the instrument may deliver biased and noisy results if
no sample preparation or correction is undertaken. In the levels of the weathering profile where ore is naturally composed of fine particles, suitable sampling associated with moisture correction may lead to reliable data. Water content can be measured from a surface moisture analyzer. In the case of very heterogeneous particles (size, morphology, chemical composition) such as earthy saprolite, the problem is far more complex and previous solutions may be ineffective. As a matter of fact, the internal pore structure and tough crystalline external surface of particles may hinder the determination of moisture content. An alternative option is likely to be explored to improve the quantification of these samples. It would be useful to develop a battery-powered, portable crusher in order to produce powder and make pellets before analysing the sample in the field. The relative waste of time for the on-site preparation of pellets would be compensated by an improvement in the quality of analytical data; thus, this analytical procedure would be of prominent interest for mining operators. This portable crusher, not available at the time of the study, appeared on the market while the present paper was being written.

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Fig. 18. Pit face measurement of Fe (◊ Not compensated, □ Water compensated).

REFERENCES


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